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**Soil Micromorphology and Site Formation Processes During the Middle to Later
Stone Ages at the Haua Fteah Cave, Cyrenaica, Libya**

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ABSTRACT

Understanding the timing and conditions of the Middle to Later Stone Age (MSA–LSA) transition in North Africa is, as with many cultural transitions during the Palaeolithic, based predominantly on archaeological and palaeoenvironmental records preserved within a small number of deep cave sediment sequences. The sedimentation processes that created these sequences, however, vary between sites, and impact on the observed cultural and environmental sequences and their interpretations. In order to use these sequences as chronological cornerstones it is vital to develop a robust understanding of the formation processes that created them. This paper utilises geoarchaeological analyses (field observations, soil micromorphology, bulk sedimentology) to examine site formation processes and stratigraphic integrity during the MSA–LSA at the Haua Fteah cave, Libya, one of North Africa's longest cultural sequences. The depositional processes identified vary in mode and energy, from aeolian deposition/reworking to mass colluvial mudflows. Identification of these changing sedimentation processes will impact greatly on the interpretation of the sequence's palaeoenvironmental and archaeological records, not least in identifying potential colluvial sediment deposition and reworking in layers identified as containing the MSA/LSA transition. This study highlights the importance of developing, through geoarchaeological analyses to cultural sequences in order to fully unravel the limitations and potential of archaeological and palaeoenvironmental records they contain.

KEY WORDS: Soil Micromorphology; Geoarchaeology; Caves; North Africa; Site Formation Processes.

INTRODUCTION

The rise and spread of Later Stone Age/Upper Palaeolithic (LSA/UP) industries within Africa ca. 50–40,000 BP (ca. 50–40 ka) marks the supposed florescence of 'fully modern behaviour' by *Homo sapiens* populations, behaviour that potentially marked increased environmental and/or social adaptations that enabled *H. sapiens* to spread across the globe (Bar-Yosef, 2002). The mechanisms and conditions of the development of LSA industries from the Middle Stone Age (MSA) industries that preceded them are subject to debate. North Africa, between Sub Saharan Africa and the Levant, is a region crucial to understanding the global spread of humans (Garcea, 2012; Foley and Lahr 1997; van Peer, 1998), yet the timing and conditions of the MSA-LSA transition in this key region still requires clarification (Barton et al., 2016).

Present understanding of the MSA/LSA transition in North Africa is, as are most Palaeolithic regional chronologies, based largely on artefacts, palaeoenvironmental proxies and dating material preserved within cave sediment sequences. Yet each of these stratigraphies was formed through processes unique to their setting and history (Farrand, 2001; Woodward & Goldberg, 2001). Changing modes and rates of deposition have been long known to impact on the taphonomy of the archaeological record through sediment removal and reworking, or changing rates of sedimentation and hiatuses (e.g. Butzer, 1971; Harris, 1989; Stein, 1987),

distorting interpretation of these cultural and environmental chronologies, and creating apparently abrupt changes in environmental proxies or technological artefact attributes, as well as 'inversions' of cultural material (Campy & Chaline, 1993; Hunt et al., 2015; Mallol et al., 2012). Therefore, to use deep cave sequences as cornerstones for regional cultural chronologies we must understand the processes that created them, and consider the archaeological records they contain in light of this understanding.

Geoarchaeological analyses are well-placed to analyse site formation processes through a range of methods and analytical scales (e.g. Bailey and Woodward 1997; Frumkin et al. 2016; Mallol et al 2008). In particular, soil micromorphology allows microscopic interrogation of sediments, enhancing field observations and complementing quantitative sediment analysis (Goldberg & Sherwood, 2006; Woodward & Goldberg, 2001), especially in studies of sedimentation rates and stratigraphic integrity beyond the reach of radiocarbon dating (e.g. Aldeias et al., 2014; Karkanas & Goldberg, 2010; Mallol et al., 2012).

This paper utilises a range of geoarchaeological techniques to examine site formation processes in the late MSA to early LSA layers at the Haua Fteah cave on the Cyrenaican coast of northeast Libya (22°3'5"E, 32°53'70"N). The Haua Fteah's cultural sequence revealed by Charles McBurney's excavations in the 1950s (McBurney, 1967) is unparalleled in North African prehistory, spanning the last glacial cycle (Douka et al., 2014; Jacobs et al., 2017), and containing cultural material from the MSA to the present (McBurney, 1967). Renewed

investigations by a multi-disciplinary team between 2007 and 2015 have combined archaeological excavation with palaeoenvironmental and chronological analyses (Barker et al., 2007, 2008, 2009, 2010, 2012; Farr et al., 2014; Rabett et al., 2014). The present study, a contribution to that project, combines soil micromorphology, bulk sedimentology, and field observations to develop a sedimentological, and taphonomic, framework for the late MSA and early LSA sediments within which to situate and interpret the existing and emerging archaeological and palaeoenvironmental records within this important cultural sequence.

The MSA and LSA in North Africa

North Africa is an important region in the appearance and spread of the LSA within, out, and potentially into Africa (Barham & Mitchell, 2008; Barton et al., 2016), at a time when late Pleistocene environmental changes had profound impacts on the climates and landscapes of the region. An MIS 5 network of Saharan rivers and lakes was exploited by populations with MSA technology (Drake & Breeze, 2016; Drake et al., 2013; Geyh & Thiedig, 2008; Osborne et al., 2008), before arid conditions in MIS4–2 – albeit with shorter humid periods during MIS 3 (Giraudi, 2005; Tjallingii et al., 2008) – forced populations to contract to the edges of the continent (Ambrose, 1998, 2003; Garcea, 2012).

Whether the LSA assemblages preserved within North African sequences represent diverse local trajectories and responses to local environments during the Late Pleistocene, or movements of populations themselves, is unclear and examination of the matter is hampered

by a lack of archaeological sites with deposits dating to MIS3 when this transition occurred (Barham & Mitchell 2008; Barton et al. 2016). The Central Sahara and Western Desert of Egypt were abandoned after MSA occupation, and remained so until the Holocene (Garcea 2004, Garcea and Giraudi 2006; Garcea 2012). In the Nile Valley, a potential refuge from Saharan aridity, deposition of artefacts of the ‘Khaterian’ complex at Nazlet Khater from 40–32ka (Vermeersch 2010 suggest a local development of LSA industries from the MSA, potentially driven by growing environmental and population pressures (van Peer & Vermeersch 2007). Likewise, whilst sites in the Maghreb may have been subject to sporadic abandonment, at Grotte des Pigeons, Taforalt, the cultural sequence appears to contain a series of stepped, local changes in technology between MSA and LSA (Barton et al 2016). The MSA /LSA transition is, however, later in the Maghreb than the Nile Valley, with the earliest LSA Iberomaurusian dated to 25-23ka (Barton et al., 2013). Situated between these two potential refuges, the MSA/LSA sequence at the Haua Fteah, where the transition from the MSA to the LSA ‘Dabban’ has been dated to 46-41 ka (Barton et al., 2015; Douka et al., 2014), occupies a central position in understanding the potential cultural connections and local trajectories of this important cultural shift.

THE HAUA FTEAH

The Haua Fteah cave is a semi-collapsed karstic cave on the northern escarpment of the Gebel Akhdar (‘Green Mountain’) limestone massif. The Gebel Akhdar (Figure 1) covers an

area ~300 x 400 km and rises in series of three escarpments to over 800 m (McBurney & Hey, 1955). Its topography and position in the path of Mediterranean Westerlies cause it to receive more rainfall than the surrounding regions. The cave is ~1 km from the present shoreline and ~50 m asl, with an opening ~80 m wide and ~20 m high. The floor of the cave consists of bare, largely dry, silty sediment which is easily mobilised by the wind, wetted in small areas by dripping from the cave roof. The floor remains largely dry during rain, yet intense storms can transport significant volumes of sediment and soil into the cave.

The massif was connected to Saharan rivers and lakes during MIS 5 (Drake & Breeze, 2016; Drake et al., 2013; Geyh & Thiedig, 2008; Osborne et al., 2008), and it probably acted as a refuge during glacial periods (Klein & Scott, 1986; McBurney & Hey, 1955; Prendergast et al., 2016; Reade et al., 2016). The Mediterranean coastline and the marine resources it contained may have increased the area's attractiveness to hominin populations, as well as providing a corridor for dispersal (Bailey & Flemming, 2008). The region's steep offshore topography means that the position of the coastline has not moved far over the last glacial cycle (Lambeck & Purcell, 2005).

The Haua Fteah's location and long cultural sequence make it pivotal to understanding prehistoric cultural change in North Africa. The 1950s excavations revealed >14 m of deposits containing cultural materials from the MSA to the historic period (McBurney, 1967). The cave was excavated in three stepped or inset trenches named by the present project, which has re-

emptied the McBurney trench of the 1955 backfill, as the Upper Trench (from the ground surface to ~2 m depth), the Middle Trench (~2–7 m, and the Deep Sounding (~7–14 m). The sediments were excavated in spits that often cross-cut sediment layers, though the relationship of the spits to stratigraphic layers was observed as the excavated proceeded downwards (McBurney, 1967). The present project has excavated 2 x 1 m trenches down the side of the Middle Trench and Deep Sounding, as well as two 30 x 30 cm sample columns from the surface to the base and excavating a further 1 m below the McBurney trench.

The earliest industry, a flake- and blade-based MSA industry classified by McBurney as 'Libyan Pre-Aurignacian', occurred in three phases within the Deep Sounding and base of the Middle Trench, with occasional lithics and shell fragments recovered throughout the Deep Sounding (Farr et al., 2014; McBurney, 1967; Rabett et al., 2014). These sediments have been dated by OSL to MIS 5, as far as the MIS 5/6 boundary (Jacobs et al., 2017). Close to the base of the Middle Trench, artefacts with 'Levalloiso-Mousterian' affinities were found by McBurney in initially high numbers from sediments for which the present project, using new ¹⁴C and OSL determinations, has provided a modelled age of 75–70 ka (Douka et al., 2014). These levels yielded two *H. sapiens* mandibles (McBurney *et al.* 1952, 1953; Tobias 1967; Trevor and Wells 1967). Artefact numbers then dropped significantly, remaining low into McBurney's Layer XXV, which he identified as containing the first LSA industry at the site, which he termed 'Dabban' (McBurney, 1967) and for which we have a modelled age of 46–41

ka (Douka et al., 2014), a date consistent with the earliest UP industries of the Levant and Europe (Benazzi et al., 2011; Higham et al., 2011; Rebollo et al., 2011). Genetic evidence suggests a migration of Levantine populations into North Africa at this time (Olivieri et al., 2006), perhaps along the Mediterranean coastline.

The nature of the MSA/LSA transition at the Haua Fteah as understood through the 10950s excavations is unclear. It occurred during a period of low artefact density, and Layer XXV was subdivided into sub-units a-e, with Levallois-Mousterian material in XXVc overlying the appearance of potentially LSA Dabban material in XXVd (McBurney 1967: 138). This potential interstratification of MSA and LSA technologies observed by McBurney may mark the presence of residual MSA traits or groups in Cyrenaica after the LSA, and/or erosion and reworking (McBurney, 1967: 138), and/or mixing of material by the spit excavation method. Similar complexity is emerging picture from the initial analysis of the small number of finds located at these depths by the new stratigraphic excavations (Farr et al., 2014; Rabett et al., 2014). The different scenarios have major implications for understanding the MSA/LSA transition, and highlight the need to situate the emerging new archaeological data from the entire sequence within a solid model of formation processes.

Sediment Sequence

Basic sediment analysis and interpretation of the sequence was undertaken during the 1950s excavations, and a broad framework for understanding the modes of sedimentation developed

(Sampson, 1967). Layers of limestone gravel and sand in the Haua Fteah sediments were interpreted as likely weathered from the cave roof (Sampson, 1967), during periods of cooling and increased physical weathering (MIS4, MIS2) and deposition of fine material was attributed to wind and water action (McBurney, 1967), in warmer phases (MIS5e, MIS3). Moyer (2003) later posited that these fine sediments were the result of inwash during pluvials (Moyer, 2003). These interpretations were, however, based solely on field observations and coarse material measurement, and in particular the deposition mechanisms of fine material remained untested.

As part of the renewed excavations, the sediment sequence was divided into five facies reflecting differences in the dominant site formation processes (Douka et al., 2014; Inglis, 2012) (Figures 3 and 4). Three fine-grained facies (Facies 1, 3 and 5) were separated by two dominated by limestone gravel and sand (Facies 2 and 4). Field observations during the new fieldwork, followed by bulk sedimentological studies, interpreted the silty layers within Facies 1 as deposited by soil inwash, events that were interleaved with heavily anthropogenically-influenced sediments (Hunt et al., 2010). Modelling of radiocarbon-dated shell fragments, and interpretation of palynological analyses and field descriptions, extended this 'sump' model into the upper part of Facies Two, where sediments were interpreted as predominantly debris flows (Hunt et al., 2015).

METHODS

This paper presents the analysis of the sediments, and the interpretation of the processes that deposited them, in the uppermost 'Levalloiso-Mousterian' MSA and 'Early Dabban' LSA layers

in the Middle Trench: Facies 3, 4 and the top part of Facies 5, which were the subject of RHI's PhD (Inglis, 2012). A multi-scalar geoarchaeological approach was employed: field descriptions were combined with high-resolution soil micromorphological analysis of each context, and bulk sedimentological analyses on the <2 mm fraction provided measurement of the fine sediment properties. In addition, two soil pits in the local *terra rossa* were dug and sampled to provide modern analogues, Pit R directly upslope of the cave opening (two samples collected) and 5m from the lip of the roof, and Pit T on the slopes below the cave, ~ 200m to the NNW of the Haua Fteah (three samples).

Field Observations After removal of the backfill and cleaning of the 1950s sections, the excavators utilised the single context system (MOLAS, 1994) to record discrete sediment layers defined by their field characteristics, e.g. colour, texture of fine and coarse material, shape, and clast orientation. Use of context divisions in sampling for sedimentological, micromorphological and palaeoenvironmental analyses ensured that data could be correlated with archaeological assemblages from each context as excavations progressed.

Soil Micromorphology Soil micromorphological sampling was undertaken of the majority of contexts on the West-Facing Profile in Facies 5–3, as well as corresponding North-Facing Profile contexts in areas of variable stratigraphy and at the depth of McBurney's Layer XXV (Figure 5). Sampling focused on upper and lower contacts of each context to characterise the mode of deposition, material within each deposit and the transitions between them. Intact

sediment blocks were removed from profiles using Kubiena tins or foil food containers, and secured with tissue paper and parcel tape to allow drying in the field before wrapping in clingfilm for transport. Blocks were air-, then oven-dried for 48 hours prior to impregnation with crystallitic polymer resin under vacuum, following capillary rise. Thirteen x 7 cm or 7 x 5 cm slices were cut from cured blocks and ground to 30µm using a Brot multiplate grinding machine. Thin-sections were examined by eye and using a Leica Wild M40 wide-view microscope (x4 to x35 magnification), and a Leitz Laborlux 12 Pol microscope (x40 to x400 magnification), under plane polarized (PPL), crossed polarized (XPL), and oblique incident light (OIL). Description followed Bullock *et al.* (1985), Courty *et al.* (1989), and Stoops (2003). Micro-fabrics within each slide were defined on the basis of changes in micromorphological characteristics within a sample, and numbered in stratigraphic order from the top of the slide, e.g. Micro-Fabric 2521:1 was the uppermost unit in the slide made from Sample 2521 (the sample number was assigned in the Haua Fteah environmental sample register), and overlaid Micro-Fabric 2521:2.

Bulk Sediment Analyses Sample columns (30 x 30 cm) were excavated through Facies 3-5 sediments on the North- and West-Facing Profiles (Figure 5). Samples at 5 cm intervals respecting context boundaries were collected for bulk sedimentological analyses, as well as palynology, phytolith analysis and tephrochronology (e.g. Barton et al., 2015; Simpson, 2014). The <2 mm fraction was analysed for: particle size distributions; percent loss on ignition

organics (%LOI organics); percent carbonate (%CaCO₃); and magnetic susceptibility. For the particle size analysis alone, carbonates were removed from the sample prior to analysis (see next section); all other bulk sedimentological analyses were carried out on the complete <2mm fraction.

Laser particle size analysis of the 0.02–2000 µm sediment fraction can distinguish a sediment's transport method or source (Folk, 1966), potentially useful at the Haua Fteah for distinguishing between colluvially- and aeolian-transported sediments. Given that the limestone from the cave left <1% residue after dissolution (Inglis, 2012), the volume of limestone required to produce this would be far greater than that dissolved in the formation of the cave. The non-carbonate fraction in the Haua Fteah sediments was therefore considered to be predominantly allochthonous. Removal of carbonates was therefore undertaken to avoid concretion of sediment particles by carbonate precipitated in the cave, and to examine the allochthonous, non-carbonate source of the material. It is noted, however, that this would also remove purely allochthonous carbonate material (if present), although it would be expected that such variation in lithology, if of a sufficient magnitude, would be visible in the micromorphological observations. The processes may also remove phytoliths from the sample, altering the silt fraction in the non-carbonate PSD. Samples were treated overnight with 10% HCl, and dispersed in 4.4% sodium pyrophosphate and distilled water. Analysis was carried out using a Malvern Mastersizer 2000, and the ultrasonic probe used throughout measurement

to ensure clay deflocculation. Samples were measured three times for 30s, and the average distribution taken. Measurement was repeated until the curve stabilised.

%LOI organics. Raised organic content of sediments may indicate anthropogenic activity (e.g. Macphail et al., 2004; Sánchez Vizcaíno & Cañabate, 1999; Stein, 1992) or palaeosol formation (Ellis & Mellor, 1995). %LOI organics was determined by heating pre-weighed, dried samples to 425°C for 24 hours and, on cooling, noting the mass lost.

%CaCO₃ in limestone cave sediments originates from multiple sources: limestone roof-spall from physical weathering (Laville, 1976), carbonate-rich source material (e.g. aeolian sources (Coudé-Gaussen & Rognon, 1993), precipitation from carbonate-rich water (White, 2007) or ash (Canti 2003). These different depositional processes highlight the need for bulk %CaCO₃ to be interpreted in the context of field and micromorphological observations. %CaCO₃ contents were determined through calcimetry following Gale and Hoare (1991).

Magnetic Susceptibility serves as a measure of ferro- or ferrimagnetic materials within a sample (Thompson & Oldfield, 1986), the concentration of which can be raised or lowered through changes in redox conditions through burning, waterlogging, wetting and drying cycles, or biological activity (Sternberg, 2001). Variation in magnetic susceptibility may also mark variation in source material, and magnetic susceptibility variability of re-deposited soils has been used in caves as a proxy for external environmental conditions (e.g. Ellwood et al., 2004), although this latter interpretation did not take into account the multiple mechanisms that may

raise magnetic susceptibility, including burning. The low frequency magnetic susceptibility of the samples was measured using a Bartington MS2B Meter.

RESULTS

Micromorphological and Field Observations

The three MSA/LSA facies are described from the base of the Middle Trench upwards. For summary micromorphological descriptions, refer to Supplementary Tables 1–5, and full descriptions, Inglis (2012).

Facies 5 (Middle Trench)- 75–65ka

Facies 5 (Contexts 528–520, West-Facing Profile; 538–521, North-Facing) in the Middle Trench consists of red and orange silt layers, some of which crossed both profiles, interspersed with combustion features and rare organic lenses. Some redder, more clayey layers lens out on the West-Facing Profile towards the back wall of the cave (e.g. 520), indicating their origin outside the cave mouth. Between Contexts 523 and 521 a series of small combustion features were spaced across the South-, East- and West-Facing Profiles. Micromorphological sampling covered Contexts 523–520 (West-Facing Profile) and 528–521 (North-Facing). Micromorphological observations (Supplementary Table 1) showed a largely consistent fine material composition of sandy silt/silt loams with micritic crystallitic to stipple-speckled fabrics. Micro-fabrics throughout Facies 5 shifted between those that contained lenses of material (fine sand, silt, or dung), and rarer micro-fabrics with more chaotic arrangements and sharp lower boundaries. These variations occurred within, and between, contexts (e.g. Context

563). Occasional crust fragments were observed on both profiles, and calcite precipitation on the West-Facing Profile (Context 521). Lenses of charcoal, ash and other charred and humified material were observed throughout. Facies 5 was capped on the West-Facing Profile by a dark red, clayey context (520) with a sharp lower boundary and stipple-speckled b-fabric (Micro-Fabrics 2000:1).

Facies 4 - 68–47ka

Facies 4 (Contexts 517–513, West-Facing Profile; Contexts 567–513, North-Facing) was characterised in the field by limestone gravel in a pale, friable orange to grey silty matrix. On the North-Facing Profile it was capped by a combustion feature (Contexts 513 and 535) that sloped down to the east. Only Context 513 continued across both profiles, grading from grey on the North-Facing Profile to orange on the West-Facing. Micromorphologically (Supplementary Table 2), the micro-fabrics in Facies 4 exhibited more distinct variation than those in Facies 4. On the West-Facing Profile, Context 517 (516 was not sampled) is a pale, limestone-dominated sandy silt loam/silt loam, with some dung lenses and a stipple speckled/micritic crystallitic b-fabric (Micro-Fabrics 2666:1 & 2). Similar micro-fabrics were observed on the North-Facing Profile, in layers of greyish, limestone-sand sandy silt loams with horizontal orientation of the coarse material (e.g. Contexts 567, 536), sometimes containing small red clayey stringers and crusts as well as calcitic hypocoatings. These micro-fabrics were interstratified with lenses of redder sandy clay loams with interleaved/sharp lower

boundaries, indicating major variation in depositional energy and character. These layers are capped by a large, weathered combustion feature (Contexts 535, 513) consisting of a series of ash and phytolith-rich layers (Inglis, 2012), potentially formed by combustion of grass-rich fuel (S. E. Jones, pers. comm. 2012).

Facies 3 - 48–34ka

Facies 3 (Contexts 509–440, West Facing Profile; Contexts 508–440 North-Facing Profile) consisted of silt layers varying in the field in compactness, colour, and texture, with dense, reddish layers of silty clay that stood proud of cleaned profiles interspersed with friable, pale yellow and orange silt layers. On the West-Facing Profile, some layers (both silty and claeey) were concreted by carbonate (e.g. Context 490). The horizontal layers largely continued across both profiles. Gravel layers became thicker and more frequent towards the overlying, gravel-dominated Facies 2. Facies 3 contained both MSA and LSA layers, separated by Layer XXV which may have contained interleaved MSA and LSA assemblages (McBurney, 1967). For ease of discussion, description of the results is divided here into these cultural groups.

The **Facies Three MSA contexts (509–498)** (Supplementary Table 3) were commonly pale orange-brown and yellow silts with little limestone gravel, with a shift towards a higher frequency of redder and more compact layers from Contexts 503–499, some of which were restricted to the West-Facing profile, and lensing out towards the back wall of the cave, indicating that they had originated from the cave mouth (e.g. Context 491). The Facies Three

MSA layers were sampled for micromorphological analysis on the West-Facing Profile only. Whilst containing less limestone sand and gravel than Facies 4, the micromorphological observations of the fine material were consistent with the underlying facies, that is, they varied between pale, sandy silt loams/silt loams containing dung lenses and stringers of fine material and other, redder, clay loams/silty clay loams with chaotically-arranged coarse material and clear lower boundaries that, in places, crossed both profiles (e.g. Contexts 508, 499, 498). A third group contained chaotically-arranged pale silty loams with micritic crystallitic to stipple-speckled b-fabrics (e.g. Contexts 506, 505). On the West-Facing Profile, all contexts from the upper part of 505 to 498 contained calcitic hypocoatings or, as in Context 503, micritic calcitic concretion of large areas of the groundmass.

McBurney's **Layer XXV** was defined from the 1950s section drawings as Contexts 497–491 in the West-Facing Profile and Contexts 470–458 in the North-Facing Profile), covering a greater depth than defined during excavation due to discrepancies on the published 1950s profile drawings. Layer XXV contains markedly red, clayey layers that alternate with paler silty layers. Those that were traceable across the two profiles were sampled on both. The micromorphology of the Layer XXV contexts (Supplementary Table 4) was dominated by silt loams/silty clay loams that contained fine material lamination, dung lenses and horizontal orientation of coarse material, although Micro-Fabric 2521:2, the lower part of Context 494/459 on the North-Facing Profile, had a chaotic arrangement towards its base. Context

493/458 had a clear lower boundary and a similarly chaotic arrangement on both the West-Facing Profile (Micro-Fabric 1014:2), and North-Facing Profile (Micro-Fabrics 2521:1, 2529:2), where it is redder. Context 491, restricted to the West-Facing Profile, was a reddish silty clay loam with a clear lower boundary, and lensed out towards the back wall of the cave, indicating it had originated in a movement from the cave entrance. Micritic calcitic void hypocoatings were present in most West-Facing Profile contexts in Layer XXV.

The **Facies 3 LSA layers** (Contexts 490–236, West-Facing Profile; Contexts 455–202, North-Facing Profile) contained an increasing intensity of gravel lenses (most prominent on the North-Facing Profile) towards the top, interrupting otherwise red-orange silty layers. The recorded stratigraphy varied between profiles, with some contexts on the West-Facing Profile divided into finer layers on the North-Facing. Contexts 490–442 (West-Facing Profile) and Contexts 547–461 (North-Facing Profile) were sampled for micromorphological analysis. The LSA layers (Supplementary Table 5) were largely pale, silty loams/silty clay loams with varying amounts of gravel and horizontal orientation to the fine and coarse material as well as dung lenses, and occasional clay stringers and crusts. These contrasted with reddish silty clay loams with sharp lower boundaries (Contexts 490, 461), some with reticulate b-fabrics (e.g. Micro-Fabric 762:1, Context 461). Micro-Fabric 754A:2 in Context 453 is unique in the Haua Fteah observations in that it contains distinct laminations and dung lenses as well as mosaic and reticulate b-fabrics, dendritic manganese staining and semi-dissolved bone associated with

neoformed minerals. On the West-Facing Profile, some contexts were heavily concreted by micritic calcite (e.g. Contexts 442, 445), and contained calcitic pedofeatures.

Interpretation of micromorphological observations

The differences in micromorphological characteristics within the Haua Fteah are subtle and the features are often undiagnostic when considered in isolation. The observed features were therefore interpreted in relation to each other and field observations. Interpretation of structure in these layers was hampered by compression by up to 4 m of overlying sediments. In addition, the site's excavation history – excavation, burial, re-excavation – may have altered existing, or produced new, redoximorphic features (e.g. dendritic manganese nodules), hampering interpretation of whether these features are related to the depositional environment of the layers or more ancient redox fluctuations.

The micro-fabrics were divided into three main groups through micromorphological observation. The first consists of pale, sandy silt to silt loams containing varying amounts of often horizontally-orientated limestone sand and gravel (Figure 6a), probably produced by roof-spalling (Farrand, 1975; Laville, 1976; Woodward and Goldberg 2001; Godberg and Sherwood 2006). They have largely stipple-specked b-fabrics, implying a lack of mechanical processes to orient clays, e.g. shrink-swelling (Kovda & Mermut, 2010), with a micritic crystallitic fine component (10-30%) (Figure 6c) interpreted as the inclusion of aeolian sediment or spalling of the cave walls. Micritic calcitic crystallitic b-fabrics may also mark

precipitation of micrite following sediment wetting and drying (Figure 6d) (Durand et al., 2010; Goldberg, 1979; Guo & Fedoroff, 1990), as well as the inclusion of ash within the sediments (Canti 2003). The nature of the micritic b-fabric in each case was therefore assessed via observation of related pedofeatures such as the presence of calcitic wood ash pseudomorphs (in the case of ash inclusion), or micritic calcitic void hypocoatings formed by percolation of carbonate-rich water (suggesting post-depositional precipitation). In Context 490 (Facies 3 Dabban) the formation of micrite and sparite crystals within the groundmass and pores suggests prolonged dripping onto, and wetting of the cave floor, extensive concreting Micro-Fabrics 2621:2-4. In Micro-fabric 2621:3 (Figure 6d), this calcitic crystal formation was so extensive that it produced a platy microstructure in a similar fashion to that formed by ice crystals in freeze-thaw sediments (e.g. van Vliet-Lanoë, Coutard *et al.* 1984; van Vliet-Lanoë 1998).

These fabrics often contained subtle, sub-millimetre-thick laminations in the fine material, mirrored in the horizontal orientation of limestone sand and gravel clasts as well as the presence of horizontal dung lenses (Figure 6e). These features were interpreted as resulting from punctuated, low-energy, aeolian deposition of fine material, producing ephemeral surfaces upon which dung was trampled, and limestone clasts fell (Goldberg, 2000). The dung, which often contains faecal spherulites (Figure 6f), is likely derived through the activity of herbivores in the cave (Brochier et al. 1992;

Canti 1998, 1999), the wide opening of provides easy access to wild animals seeking shelter. These surfaces may have been sporadically wetted, producing surface crusts (Goldberg, 2000; Valentin, 1991), or subject to small-scale washes of clay-rich material from inside the karstic system or through the cave mouth, leaving depositional crusts (Figure 7a) (Bresson & Valentin, 1994; Pagliai & Stoops, 2010; West et al., 1990). Fragments of both sets of crusts were observed, indicating trampling or bioturbation post-formation. In addition, iron-impregnated aggregates within the sediments with a groundmass, internal structure and b-fabric different from the surrounding material (Figure 7b) were interpreted as fragments of soil, 'pedorelicts', trampled into the cave by animals or people (Macphail and McAvoy 2008, Boschian 1997; Goldberg 1979a). Micro-fabrics within this first group of features were therefore interpreted as the product of aeolian deposition and reworking on a 'dusty', mainly dry, cave floor, similar to the modern cave floor.

The second main group of deposits examined were silty clay to silty clay loams, generally reddish in colour (difficult to assess where slides vary in thickness). Their b-fabrics were largely stipple-speckled with some mosaic-speckling (Figure 6b), the latter indicating a limited impact of shrink-swell processes, such as drying of a saturated sediment (Cremeens, 2005), and were thus interpreted as wet movements of clayey, potentially soil, material. Some contained a partially micritic crystallitic b-fabric, usually related to calcitic infillings and hypocoatings and therefore interpreted as post-depositional calcite precipitation. Coarse material was often

arranged chaotically, consistent with a mass depositional event, and the lower boundaries of these micro-fabrics are often clear, sharp (Figure 7c) and therefore potentially erosive. On occasion, silty clay coatings to voids were present in micro-fabrics directly below these layers (Figure 7d), marking the drainage of water through the profile (French et al., 2009); their depth restriction suggests that they were linked to short periods of small-scale illuviation, consistent with clayey water draining from a slurry. These micro-fabrics, and their similarity to the reddish, clayey local soils were interpreted as inwash of soil material through the cave mouth or elsewhere via the karstic system.

Not all observed micro-fabrics fitted perfectly into these two groups. A third group of micro-fabrics consisted of silty fine material, coarse limestone sand and partially micritic crystallitic b-fabric similar to the 'dusty' micro-fabrics, yet their chaotic arrangement and sharp lower boundaries suggest deposition in a mass movement, indicating that these layers represented reworking of cave-floor material by mass movements. Micro-Fabric 754A:2 (Context 453) contained a unique combination of features, including distinct horizontal laminations and dung lenses, with a mosaic/reticulate b-fabric, dendritic manganese staining and partially-dissolved bone (Figure 7e, f). These suggested a layer formed through aeolian deposition and reworking that had been subject to prolonged wetting, perhaps marking repeated dripping in this area from the cave roof.

Anthropogenic and biogenic impacts on the sediments were observed

micromorphologically, with combustion features in Facies 4 and 5 containing layers of finely commuted charcoal and ash (Figures 8a, b). Rare charcoal fragments, bone splinters and ash lenses were observed throughout, marking a continually reworked 'background' of activity debris. Vesicular silica aggregates observed micromorphologically (Figure 8c) mirrored 'slag' fragments recovered in the 1950s (McBurney, 1967), formed from the melting of a silica-rich fuel such as grasses (Canti, 2003; Macphail & Cruise, 2001). Distinct from rounded, potentially geogenic clasts, angular splinters of chert were interpreted as knapping micro-debitage (Figure 8d, Angelucci, 2010). In the large feature from 513/535, ash layers containing calcitic wood ash pseudomorphs (Figure 8e) indicated the use of wood as fuel (Canti 2003), and the presence of calcitic hypocoatings (Figure 8f) within these ash layers suggest that the feature had been subject to post-depositional wetting and weathering, indicating the feature's prolonged exposure on a surface at the top of Facies 4.

Bulk Sedimentology

Particle size analysis

The particle size distributions of the non-carbonate <2mm fractions of the Haua Fteah sediments were largely consistent (Figure 9). Whilst there was variation in mode particle sizes, the distributions were all bimodal, with a clay peak around 0.17 μ m and a 15.63–44.19 μ m silt peak. The soil pit distribution (Pit T) shared this bimodal distribution with the Haua Fteah

samples, and, bar a $\sim 353\mu\text{m}$ sand peak, the limestone residue lay within the size ranges of the Haua Fteah samples.

The consistency in particle size distributions in the non-carbonate fraction indicates that this fraction of the Haua Fteah sediments did not vary extensively in source or transport mode throughout Facies 5–3. Similarity between the Haua Fteah samples and the soil pits suggest that the local soils, formed on the local limestone, were the primary source of the non-carbonate sediment. The more marked coarse silt peak in the particle size distribution in the soil pit, when compared to the limestone residue, indicates another potential input to the soil, and, given the peak's sorting, may represent a far-field input of aeolian material, common in Mediterranean soils (Muhs et al., 2010; Yaalon, 1997). It is possible that the similarity between the limestone residue and the non-carbonate fraction may be influenced by the dissolution of limestone sand within the sample producing non-carbonate residue, yet as mentioned earlier, the very low non-carbonate content of the limestone (<1%) suggests this addition would not be enough to skew the PSD. The removal of carbonate from the samples prior to measurement does, however, raise the possibility that a carbonate-dominated aeolian input directly into cave has been removed in this analysis.

% Loss On Ignition

All the Facies 5–3 %LOI values from the Haua Fteah were at or below those of the modern soils (Table 1), suggesting a relative reduction in the organic content of the material after

deposition, or a lower initial organic content to the source material. The values were very consistent (Standard deviation = 0.5%), rising slowly towards the upper part of the sequence (Figure 10), with only two small distinct peaks on the North-Facing Profile, one corresponding to combustion features (Contexts 568 and 564) and a context containing frequent dung lenses (Micro-Fabric 2058:4, Context 536). Such low values and narrow standard deviation means that it is impossible to infer that these peaks are meaningful. These contexts, which from micromorphological observations should have contained significant amounts of organic material, had values similar to the modern soils, indicating that in sediments of this age (>30,000 years old) %LOI values have, unsurprisingly, decayed to the point at which they are no longer meaningful, and little stock should be placed on their interpretation.

%CaCO₃

The Haua Fteah samples contained markedly higher %CaCO₃ values than the soil (Table 1), indicating the addition of carbonate to the sediments if, as appears from the PSDs and field observations, that they are dominated by reworked soil material. This addition is likely the result of cave wall weathering and post-depositional carbonate precipitation (Goldberg & Sherwood, 2006; White, 2007), although it is possible that an aeolian carbonate component may also be contributing to the %CaCO₃. Fine clastic carbonate material may also have been transported into the shelter in mudflow events. Micromorphological observations of ash, which may also have contributed carbonate (Canti 2003) to the sediments are restricted to the thin

lenses of anthropogenic material in Facies 4 and 5. Broadly, the largest %CaCO₃ peaks corresponded to observations of frequent/common coarse sand in the late Facies 5/4 sediments (Contexts 520–516, West-Facing Profile; Contexts 537–536, North-Facing Profile, Figure 10), yet West-Facing Profile contexts cemented with calcite also corresponded to %CaCO₃ peaks (Contexts 504–503, 445–442), indicating an equifinality irresolvable without field or micromorphological data.

Magnetic Susceptibility

Magnetic Susceptibility values varied widely between the soil pits (mean values of 365.45 and 584.8 m³kg⁻¹), yet were still largely higher than the cave values (Table 1). Peaks in magnetic susceptibility corresponded largely to contexts identified micromorphologically as inwashed soil material (e.g. Contexts 508, 498, 493/458–490, Figure 10), yet others were associated with 'dusty' cave floor environments (e.g. Contexts 521–523, 563), suggesting that another mechanism had raised the values: anthropogenic influence on the sediments, or, less likely in caves where pedogenic process are weaker, weathering of the sediments. This equifinality meant that the nature of the large jumps in magnetic susceptibility at the top of the sampled area, beyond micromorphological sampling, remain ambiguous. Given field observations of charcoal in Context 441, these peaks may have been related to burning.

A negative correlation (-0.588) between %CaCO₃ and magnetic susceptibility was illustrated in the major peaks of magnetic susceptibility being largely accompanied by low

%CaCO₃ values and vice versa. This, and the relatively lower magnetic susceptibility of most of the Haua Fteah sediments compared to the soil, may result from the magnetic susceptibility of consistent soil material in the cave being 'diluted' by variable amounts of diamagnetic carbonate sand and silt (Dearing et al., 1985) from the cave walls. Yet correlation of the variables was only moderately negative, and an R² value of 0.35 indicates a broad spread of the data - the impact of the coarse fraction can, in future be assessed by restricting measurements to the <83um fraction (Woodward, 1997a, b). It is also possible that magnetic susceptibility variability may have been influenced by source material variation (Ellwood et al., 1997) such as incorporation of allochthonous carbonate material, or post-deposition burning or weathering (Tite & Mullins, 1971).

DISCUSSION

Integrating Field Observations, Micromorphology and Bulk Sedimentology

The effectiveness of particle size analysis to distinguish between aeolian and colluvial deposition appears to have been hampered by the apparent local sediment source, as the variation between aeolian-dominated and colluvial/inwash deposition observed in the micromorphology was not identifiable in the particle size distributions. Removal of the carbonate fraction may have removed a well-sorted aeolian carbonate component, yet it is unlikely that this component would be solely carbonate - a non-carbonate aeolian element would be expected to remain identifiable. A lack of a far-field aeolian signature in the Haua

Fteah would be unsurprising given its position on the northern side of the Gebel Akhdar facing away from the Sahara. The lack of a large exposed continental shelf during periods of low sea level also limited another past source of aeolian sediment (Lambeck & Purcell, 2005). In addition, the fine material composition observed micromorphologically is relatively consistent through the sequence (though more clayey within the 'mudflow' group); marked variations in fine material source indicating pulses of (for example) carbonate beach dune sands would be expected to be visible in the thin sections, and are not. The lack of variability seen in the non-carbonate PSDs between layers with quite different micromorphologically-observed depositional histories indicate therefore that whilst this data can therefore inform on the local source of the non-carbonate fraction, micromorphological observation is required to understand the process of deposition.

Small peaks in bulk %LOI organics appeared to correlate with field and micromorphological observations of charred material, and dung (536), yet given the low values and restricted variation of the dataset, the %LOI values can add little confidently to the interpretation of the sequence; the variation may even be influenced more by variation in sediment lithology than organic content (Santisetban et al. 2004).

Magnetic susceptibility peaks corresponded frequently with layers interpreted through micromorphology and field observations as soil inwash (e.g. Contexts 498 and 508), rather than parts of the stratigraphy that had been burnt. An exception may have occurred at the top of

Facies 3 (Context 441), where charcoal observed in the profiles and an increase in occupation material (McBurney, 1967) may account for large magnetic susceptibility peaks. In addition, not all contexts interpreted micromorphologically as colluvial/wash events were accompanied by peaks; contexts interpreted micromorphologically as colluvial mass reworking of cave-floor sediment (e.g. Contexts 497 and 563) lacked high magnetic susceptibility values, confirming that magnetic susceptibility may inform on sediment source, but not depositional context.

Given the range of processes identified through field observations and micromorphology that added carbonate to the sediments, it is largely impossible to interpret the bulk %CaCO₃ data in isolation. Broad peaks in %CaCO₃ corresponded to increased limestone sand and gravel in Facies 4, whilst some of the lowest %CaCO₃ readings corresponded to contexts with high magnetic susceptibility interpreted as soil inwash (e.g. Contexts 490 and 498). Yet other %CaCO₃ peaks marked contexts containing secondary calcite precipitation (Contexts 503 and 504) and little limestone sand. No distinct peaks accompanied calcitic ash deposition linked to anthropogenic activity in Facies 5, although this may have been due to the small quantities of ash involved.

It is clear that the bulk sedimentological parameters measured here are subject to issues of equifinality. Further clarity of the composition of the bulk sediments through the investigation of variation in sedimentological characteristics between different sediment fractions (e.g. that of the <63µm fraction following Woodward and Bailey 2000) or sediment

sourcing using SEM, ARD or FTIR could, in future, be carried out to understand more fully the sources of the material and properties measured. Micromorphological analysis of the Haul Fteah sediments, however, appears the most robust method of analysis and interpretation of the final depositional processes that ultimately shaped its stratigraphy.

Site Formation Processes at the Haul Fteah in Their Mediterranean Context

The Haul Fteah sediments show repeated shifts between fine material deposition dominated by dry, 'dusty' conditions and sporadically wetted surfaces, sometimes with limestone clast deposition from the cave walls, and sporadically interrupted by wet, colluvial mass movements. These processes are expressed in numerous other Mediterranean Quaternary cave sequences (Frumkin et al., 2016; Woodward & Goldberg, 2001), many of which contain archaeological sequences key to understanding population dynamics and change in the region.

Mass colluvial deposition of soil material is a common sedimentation noted in Mediterranean caves (Albert et al., 1999; Aldeias et al., 2014; Bar-Yosef et al., 1992; Boscian, 1997; Frumkin et al., 2016; Goder-Goldberger et al., 2012; Goldberg & Bar-Yosef, 1998; Woodward & Goldberg, 2001). Whether driven by climate change and/or human impact, these movements mark landscape destabilisation (Frumkin et al., 2016; Wainwright, 2009), and are largely recorded in caves opening in shallow inclines, or those with chimneys such as Konispol, Albania (Schuldenrein, 1998, 2001) and Tabun, Israel (Albert et al., 1999). At the Haul Fteah these processes did not dominate the Facies 5–3 sediments, but instead

occasionally punctuated an otherwise dry cave floor environment. Whilst this reflects the cave's susceptibility to collecting colluvially-deposited sediments, this process was only one of a small number of sedimentation processes that also included aeolian deposition/reworking and roof spalling.

The aeolian nature of the 'dusty' deposits likewise have Mediterranean parallels. Whilst the Haua Fteah lacks wind-deposited beach sand in contrast to caves adjacent to shorelines such as Vanguard and Gorham's Caves, Gibraltar (Macphail et al., 2000), (e.g. in Vanguard and Gorham's Caves, Gibraltar Macphail et al., 2000), fine-grained, aeolian sediments were reported from Akrotiri Aetokremnos, Cyprus (Mandel & Simmons, 1997), Khef el Ahmmar, Morocco, (Barton et al., 2005), Abri Pataud, France (Farrand, 1975), and Klithi, Greece (Woodward 1997a) variously attributed to local sources such as floodplain sediments or, as in the Haua Fteah, surrounding hillsides.

The physical weathering of limestone from cave walls has long been documented in the Mediterranean (Collcutt, 1979; Farrand, 1975; Laville et al., 1980), with micromorphological attributes suggesting freeze-thawing as the driver of physical weathering at Theopetra, Greece (Karkanis, 1999, 2001), Abric Romani, Spain, and Grotte des Pigeons, Taforalt, Morocco (Courty & Vallverdu, 2001). Elsewhere, for example at Ksar Akil, Lebanon and Franchthi Cave, Greece (Farrand, 2001b), independent proxies suggest that the areas did not experience freezing temperatures and that chemical weathering or wetting/drying cycles drove weathering

(Goldberg & Sherwood, 2006). Micromorphological structures produced by freeze-thawing (e.g. lenticular structures, van Vliet-Lanoë, 1998) were not observed in the Haua Fteah, yet its largely dry cave floor may have inhibited their formation. There is often snow on the Gebel Akhdar in winter and current winter temperatures in Benghazi can reach 0°C (<http://www.climatetemp.info/libya/benghazi.html>). Regional cooling during glacials and stadials would have further lowered temperatures, suggesting that freeze-thaw could have contributed to physical weathering at the Haua Fteah during certain periods or seasons in addition to ongoing chemical weathering.

The sediments investigated here show limited carbonate concretion or flowstone development on the West- and North-Facing Profiles, unlike caves such as Qesem Cave (Gopher et al., 2010; Karkanas et al., 2007) or Emanuel Cave (Goder-Goldberger et al., 2012), Israel. The South-Facing Profile sediments, however, and the adjacent northernmost end of the West-Facing Profile, are heavily concreted, probably linked to their proximity to the modern dripline. The timing of calcite deposition remains unclear, and may considerably post-date the deposition of the units. This spatial variability suggests that even within a few metres, the impact of depositional processes can vary, and is dependent on cave morphology as well as environment.

The processes that deposited the Haua Fteah sediments are observed in caves across the Mediterranean. Whilst the dominant processes vary from site to site based on cave morphology

and setting, the variability in sediment energy and rate of deposition between the processes highlight the necessity of understanding site formation processes at each and every site in order to understand the archaeology and palaeoenvironmental sequences within them.

The Haua Fteah Sediment Sequence

Whilst the detailed consideration of the environmental drivers of sedimentation at the Haua Fteah with regards to Late Quaternary regional environmental change is beyond the scope of this paper, there are general trends that show some broad trends and correlation with environmental (Douka et al, 2014; Inglis 2012). The Facies 5 sediments in the Middle Trench, with a modelled age of deposition between 75–64 ka (Douka et al., 2014), correspond to MIS5a and the beginning of MIS4, and consist predominantly of aeolian-deposited fine material, interrupted by reworking of cave-floor material in mudflows indicating occasional landscape instability. There is limited physical weathering of the cave walls shown in occasional limestone clasts, yet the sediments are dominated by fine material, indicating that rarely reached the lows necessary to accelerate physical weathering of the cave walls through freeze-thaw processes. The combustion features appear to have been deposited on dry, unconsolidated surfaces, consistent with the overall environment of aeolian deposition and reworking.

The Facies 4 sediments, with a modelled age of deposition between 68–47 ka, (corresponding to MIS4 and the start of MIS3), are dominated by limestone gravel and sand,

the result of increased physical weathering, and reflected in higher %CaCO₃ values. Depositional environments remained similar to those of Facies 5, but with increased roof-spalling along with fine aeolian deposition and reworking, likely the result of MIS4 driven cooling activating freeze-thaw weathering. Sporadic wetting of the largely dry sediments continued, but more numerous mudflows into the cave (compared with Facies 5), marked increased frequency of landscape instability. At the top of the facies, extended exposure and weathering of a large burning event supports a sediment hiatus suggested by the chronology (Douka et al., 2014; Inglis 2012).

Facies 3 sediments, with a modelled age of between 48–34 ka (equivalent broadly to MIS3), were marked by smaller quantities of carbonate gravel and sand indicating an overall reduction in physical weathering, and therefore potentially warmer temperatures than in MIS4. Limestone clast deposition in contexts towards the top of the facies indicate shorter-lived drops in temperatures. As in Facies 4, the deposition of fine material varied between ‘dusty’ environments and mudflow deposition. The red, clayey, mudflow sediments with high magnetic susceptibility towards the top of the facies, accompanying the layers containing limestone clasts, suggest the increasingly frequent transport into the cave of well-developed soils or subsoil horizons from the landscape during periods of landscape instability, potentially driven by increasingly cool periods along a downward temperature trend towards the gravelly-dominated MIS2.

Whilst the reconstruction of the environmental drivers behind the changing sedimentation processes at the Haua Fteah require more discussion than is possible here, this discussion demonstrates that the environmental conditions in which caves were inhabited had profound implications for the nature of the record that is preserved of these occupations.

SITE FORMATION PROCESSES AND THE HAUA FTEAH STRATIGRAPHY

The analysis of the Haua Fteah sediments has identified two predominant modes of deposition (fine aeolian deposition/reworking, and mudflows of external or internal material) that represent very different taphonomic pathways, and major interpretative implications, for the archaeological and palaeoenvironmental material preserved within them.

In terms of the impact on the interpretation of palaeoenvironmental proxies, ‘dusty’ deposition, with its continual small-scale surface reworking of fine material on a dry cave floor, would, with its slow net rate of deposition, allow sediments to blow into and around the cave over an extended period of time, mixing with material already deposited in the cave. These layers likely contain assemblages of environmental proxies that have been subject to ‘time averaging’; for example, rapid vegetation change may not be recorded if the mode of deposition in the cave remained constant; new proxy assemblages would be mixed with those already within the cave, creating assemblages that do not bear direct relation to the outside environment, and which masked shorter-term environmental fluctuations. In contrast, the rapid deposition of the mudflow events would avoid this degree of time averaging, but also may

have transported and re-deposited proxy archives inherited from soils or sediments already existing in the landscape. Abrupt changes identified in palaeoenvironmental proxies as recorded in the CPP sample columns may therefore be more due to changes in taphonomic pathway than environment, whilst periods of consistency may mask a more complex picture, highlighting the need to ground proxy interpretation within the detailed stratigraphic framework that this study has provided.

The varying rates and modes of deposition would have similar effects on the archaeological assemblages within them. Low energy, aeolian depositional environments and their slow net rate of aggradation meant that artefacts were exposed on, or very near to, the surface for variable periods of time, when they may have been trampled by animals and/or humans. In addition, during their exposure, artefacts may have been moved, recycled or completely removed (Bailey, 2007; Vaquero et al., 2012). Material on these slowly aggrading surfaces may have been deposited in a single episode of activity, or through multiple episodes during low net sedimentation, thus forming palimpsests of increasing time depth (Bailey & Galanidou, 2009; Stern, 2008). These 'dusty' layers therefore have relatively low chronological and behavioural resolution, even if the artefacts have not been moved by the sedimentation processes or later activity, and may even contain artefacts from different cultural groups, blurring cultural transitions that were in fact quite abrupt. In contrast, the mudflow layers were deposited in single, or a series of very temporally constrained depositional event (e.g. days),

potentially burying artefacts soon after deposition, but also potentially producing archaeologically 'sterile' layers that could appear to represent occupation hiatuses, but may only mark a day of rapid deposition. Again, the taphonomic impact of the sediments must be borne in mind when interpreting the changing densities of archaeology throughout the sequence, and their behavioural implications.

These changing processes, and their changing depositional energies, are crucial to understanding the stratigraphic integrity of the MSA/LSA 'transition' layer, Layer XXV. As noted previously, McBurney identified reworking, or population turnover, as a potential reason for the apparent interstratification of Levallois-Mousterian and Dabban material in Layer XXV (McBurney 1967). The identification of mudflow deposits within Layer XXV provides support to the reworking hypothesis, yet Layer XXV also contain layers built up through dusty/aeolian sedimentation – if the different artefact assemblages were contained within these layers, they may not have been reworked or mixed by mudflows. Correlation of the McBurney archaeological archive with specific CPP contexts, the degree of resolution required to resolve the question of the integrity of the transitional layers, is impossible given its excavation history. The renewed CPP investigations provide the increased stratigraphic resolution which could address these issues, yet the Facies Three layers in Trench M have yielded a very low number of artefacts (Farr et al. 2014). We are left with, on one hand, a poorly stratigraphically constrained archive with enough artefacts (still low in number – McBurney 1967) to identify

shifts in technology, or on the other, a well stratigraphically-constrained record which contains few artefacts, and none that are able to robustly pinpoint where in the stratigraphy this transition occurs. An enhanced understanding of the taphonomy of the key MSA/LSA layers, and its implications for the archaeological record within them, may therefore remain out of reach.

CONCLUSION

Detailed geoarchaeological analysis of the Haua Fteah sediments integrating field observations, micromorphology and bulk sediment analyses has demonstrated the modes of formation of the sediments comprising the late MSA/early LSA layers of one of the key North African cultural sequences. Sedimentation throughout this period, characterised as Facies 5–3, was dominated by low-energy, slow net aeolian sedimentation and reworking, interrupted by mass movements of material from outside or inside the cave. These changing depositional environments have important implications for understanding the archaeological sequence and the cultural transitions within it, because the layers represent different taphonomic histories.

The varying site formation processes observed through the MSA and LSA levels, the limited archaeological resolution of the sequence, and the low numbers of stratigraphically-constrained artefacts from the new excavations, mean that a higher-resolution assessment of the MSA/LSA transition and its palaeoenvironmental conditions through the Haua Fteah sequence remains difficult. Full analysis of the findings from the new excavations will improve

the stratigraphic understanding of the newly-recovered artefacts, their sedimentological context, and the extent to which reworking may have affected the observed sequence, though it remains possible that, given issues of sediment redeposition, the Haua Fteah sequence may not preserve a high-resolution archive of the nature, conditions and timing of some of the key cultural transitions that took place within the ca. 140,000-year human frequentation of the site. Most importantly, the sediment formation processes recognised in the Haua Fteah, and the issues of archaeological interpretation raised by these findings, are consistent with those of many other human occupation caves in the Mediterranean region (and indeed beyond), highlighting the necessity to employ a battery of geoarchaeological analyses to establish the limitations and potential of the archaeological and palaeoenvironmental records contained within them, and the regional and global chronologies they underpin.

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FIGURES

Figure 1: (a) Location of the TRANS-NAP study area and the Haua Fteah. Drawing by D. Kemp; (b) Looking East along the northern escarpment of the Gebel Akhdar towards the Haua Fteah. Photo: R. Inglis. (c) View into the Haua Fteah shelter from North. Figure circled for scale. Laser scans of the inside of the cave showing (d) an aerial view and (e) cross-section looking East. Scans: J. Meneely and B. Smith.

Figure 2: Photo of the Middle Trench North-Facing Profile, showing the sedimentological facies distinguished and discussed in this paper. Sediments vary between limestone gravel-dominated facies (Facies 2 and 4) and those with fine, silty sediment (Facies Three and Five). Note large burning feature extending across the section from the left at the top of Facies Four. Holes from the removal of the first phase of micromorphological sampling are visible ~2.5m from the base of the trench (ranging pole divisions are 50 cm), whilst the holes to the left of the section were made through removal of samples for OSL dating. Photo: G. Barker.

Figure 3: Summary of field facies, existing cultural divisions from McBurney (1967), and dates from Douka et al. (2104).

Figure 4: Profile drawings of the West and North-Facing profiles showing locations of samples discussed in this paper. Large grey rectangles show location of bulk sedimentological sample columns, and black rectangles mark individual micromorphology samples. Section drawings: L. Farr.

Figure 5: Schematic diagram of Facies Five–Three West- and North-Facing Profiles showing relationship between the interpreted site formation processes and bulk sedimentological variables. Where they correspond to micromorphological samples (grey squares), the sediment interpretations are based on micromorphological observations discussed in the text, and where they were not sampled for micromorphological analyses, are based on field observations. Grey lines on bulk sedimentology graphs show mean soil pit values. For colours see online version.

Figure 6: Photomicrographs of key features in the Haua Fteah sediments I: (a) limestone sand and gravel interpreted as roof spall, in silty clay material, PPL (Micro-Fabric 961:2; Context 504; Facies 3 MSA); (b) Stipple- to mosaic-speckled b-fabric, indicating lack of shrink-swelling processes, XPL (Micro-Fabric 939:2; Context 508; Facies 3 MSA); (c) Stipple speckled to micritic crystallitic b-fabric – the micritic calcite may be linked to aeolian deposition, roof spalling, or post-depositional precipitation. XPL (Micro-Fabric 2028:2; Context 521; Facies 5); (d) Micritic and sparitic calcitic precipitation linked to persistent wetting of the sediments, the formation of crystals leading to the development of a platy structure, XPL (Micro-Fabrics 2621:2–4; Context 490; Facies 3 Dabban); (e) dung lenses and fine mineral material laminations marking ephemeral surfaces, PPL Micro-Fabric 2021:2;

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TABLES

Table 1: Summary statistics for the soil pit and Haua Fteah Facies 5–3 bulk sedimentological characteristics.

Figure 1: (a) Location of the TRANS-NAP study area and the Haua Fteah. Drawing by D. Kemp; (b) Looking East along the northern escarpment of the Jebel Akhdar towards the Haua Fteah. Photo: R. Inglis. (c) View into the Haua Fteah shelter from North. Figure circled for scale. Laser scans of the inside of the cave showing (d) an aerial view and (e) cross-section looking East. Scans: J. Meneely and B. Smith.

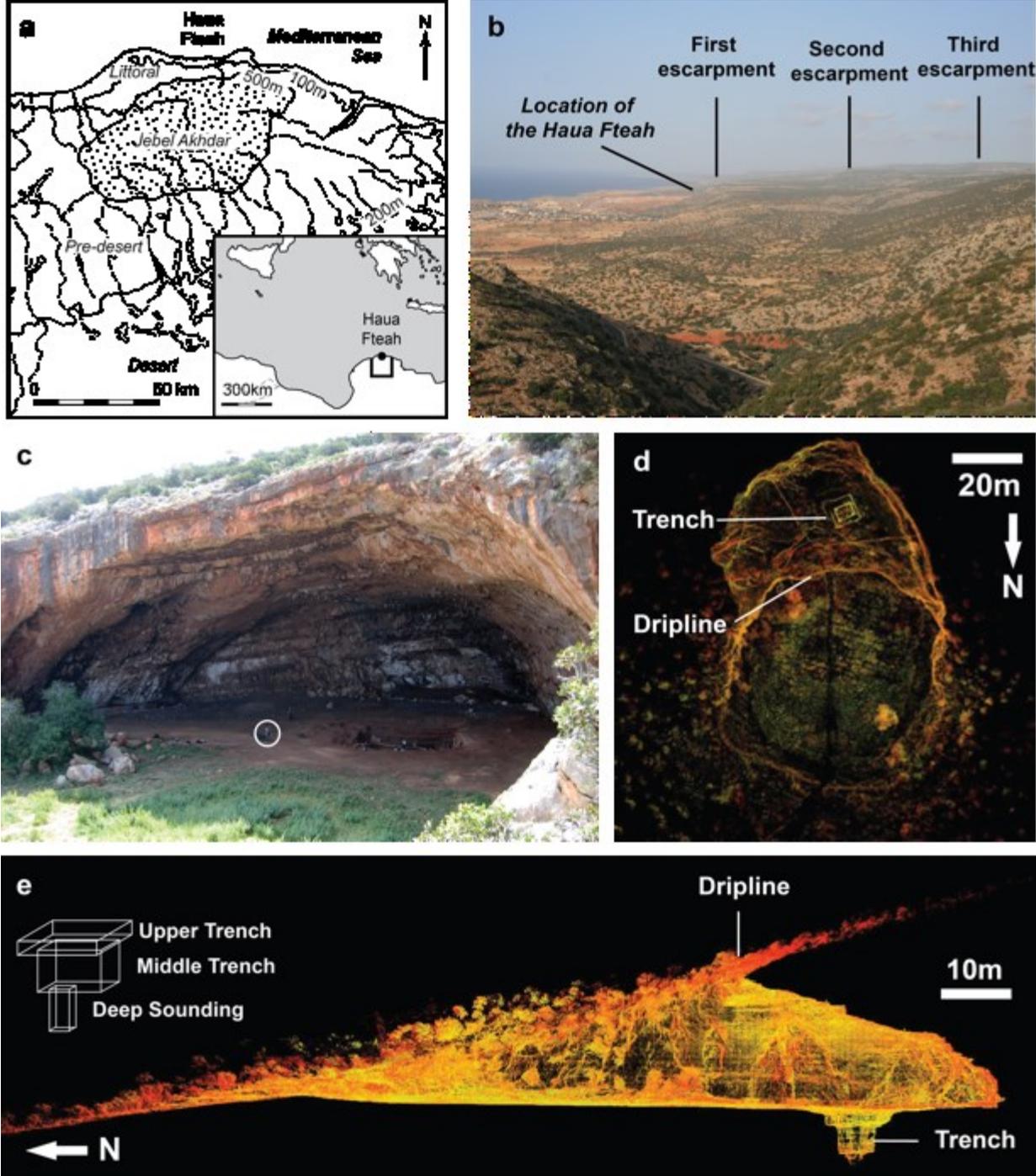


Figure 2: Photo of the Middle Trench North-Facing Profile, showing the sedimentological facies distinguished and discussed in this paper. Sediments vary between limestone gravel-dominated facies (Facies 2 and 4) and those with fine, silty sediment (Facies Three and Five). Note large burning feature extending across the section from the left at the top of Facies Four. Holes from the removal of the first phase of micromorphological sampling are visible ~2.5m from the base of the trench (ranging pole divisions are 50 cm), whilst the holes to the left of the section were made through removal of samples for OSL dating. Photo: G. Barker.

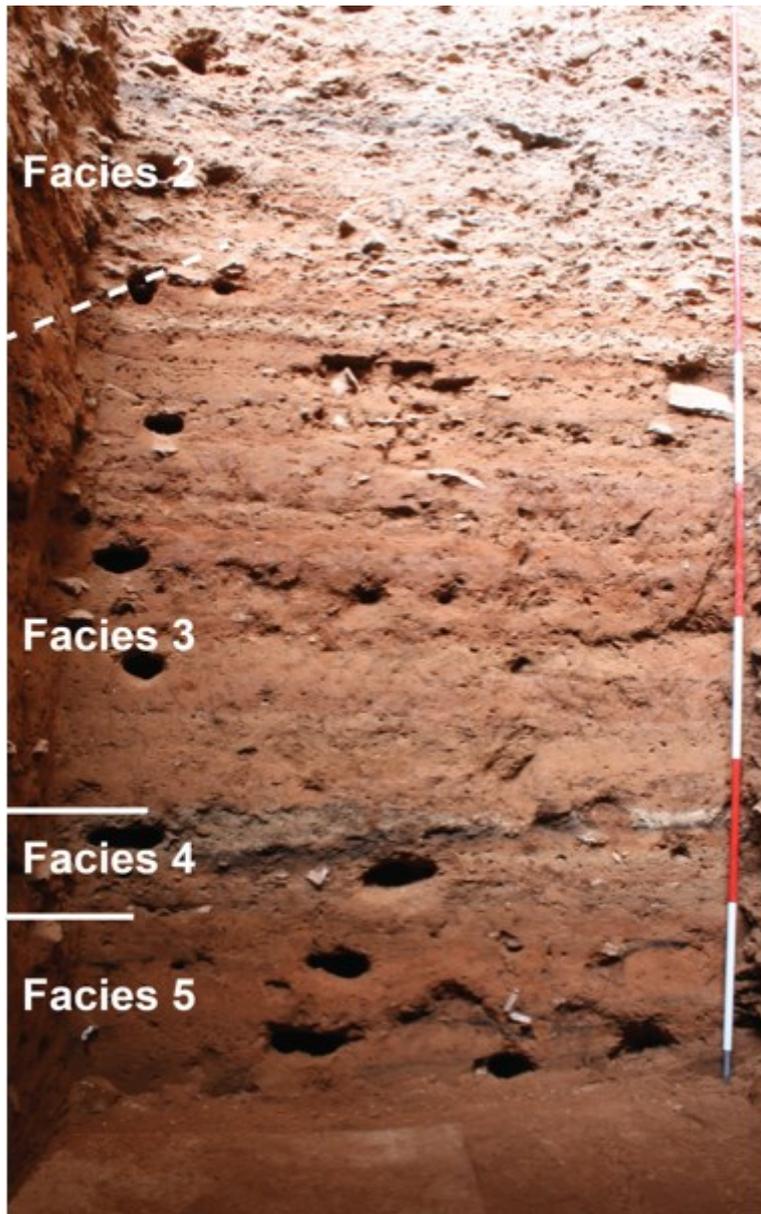


Figure 3: Summary of field facies, existing cultural divisions from McBurney (1967), and dates from Douka et al. (2104).

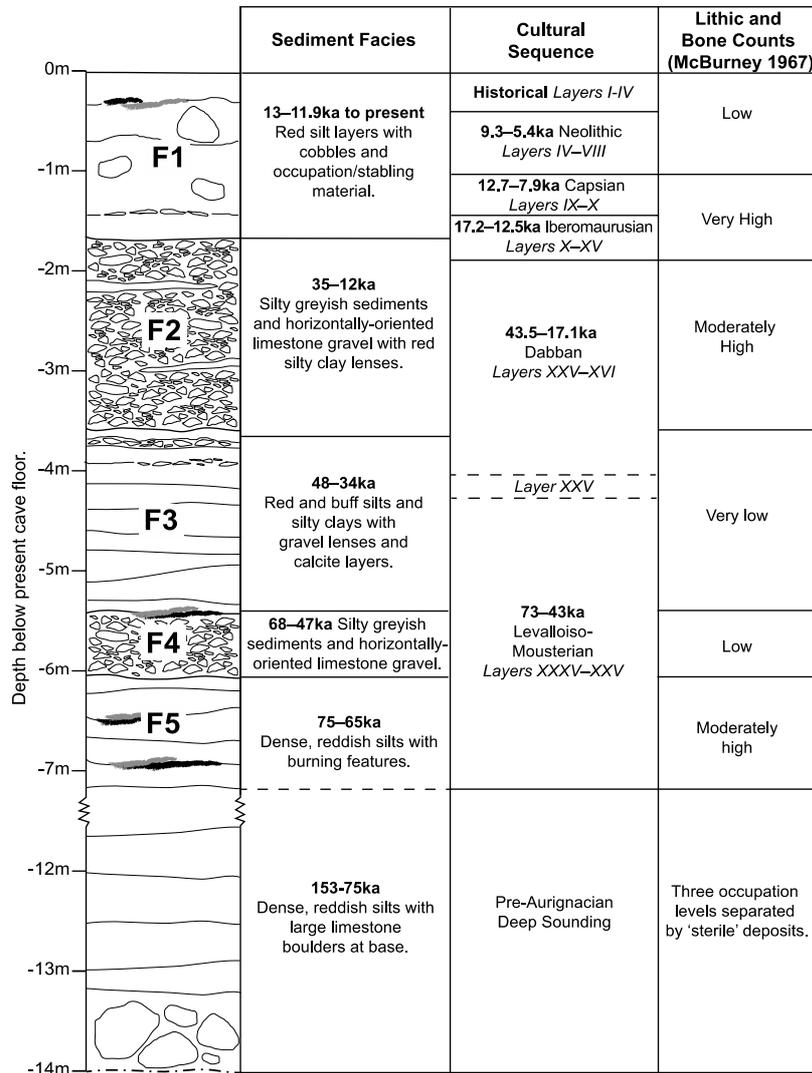


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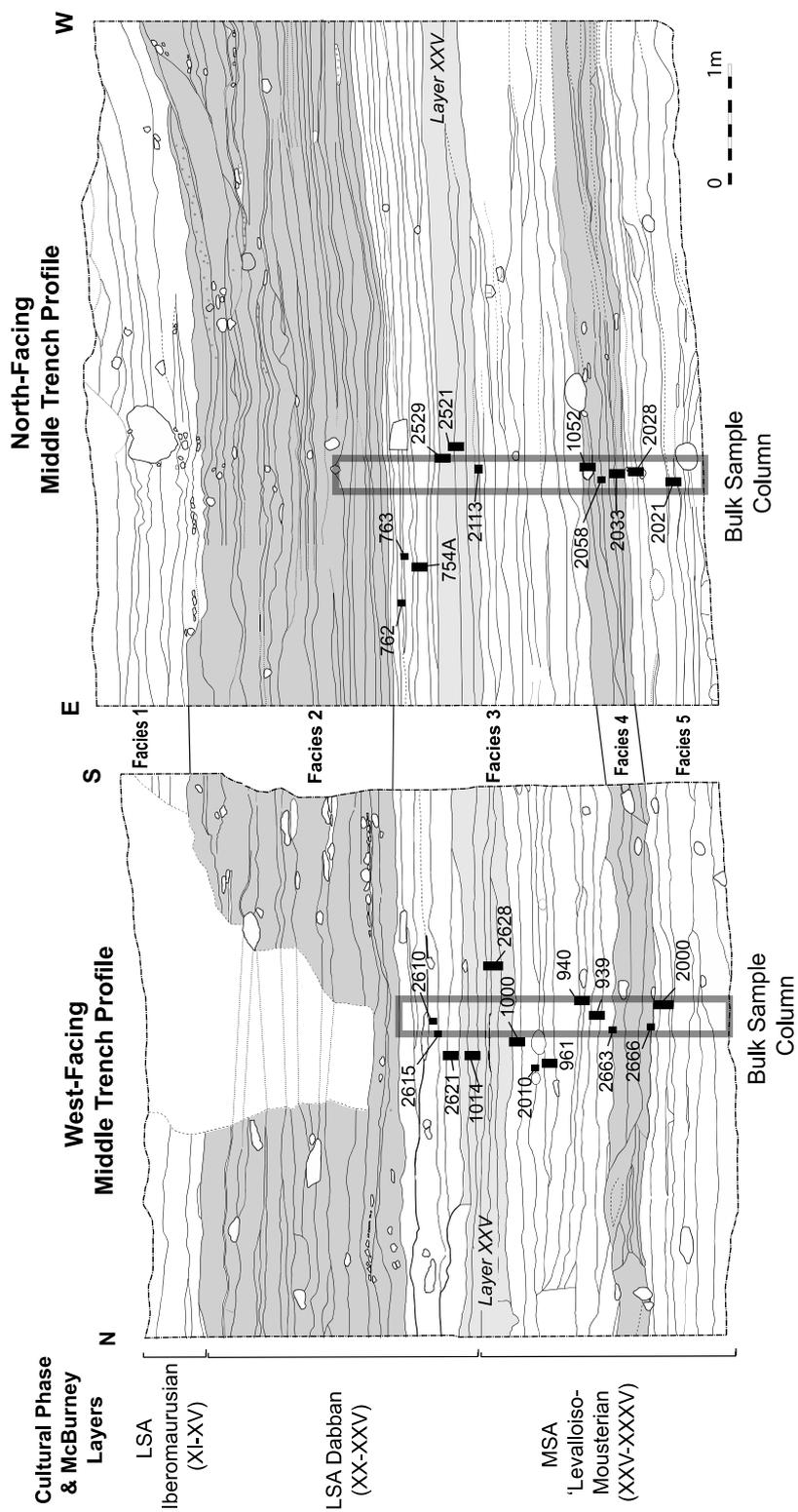


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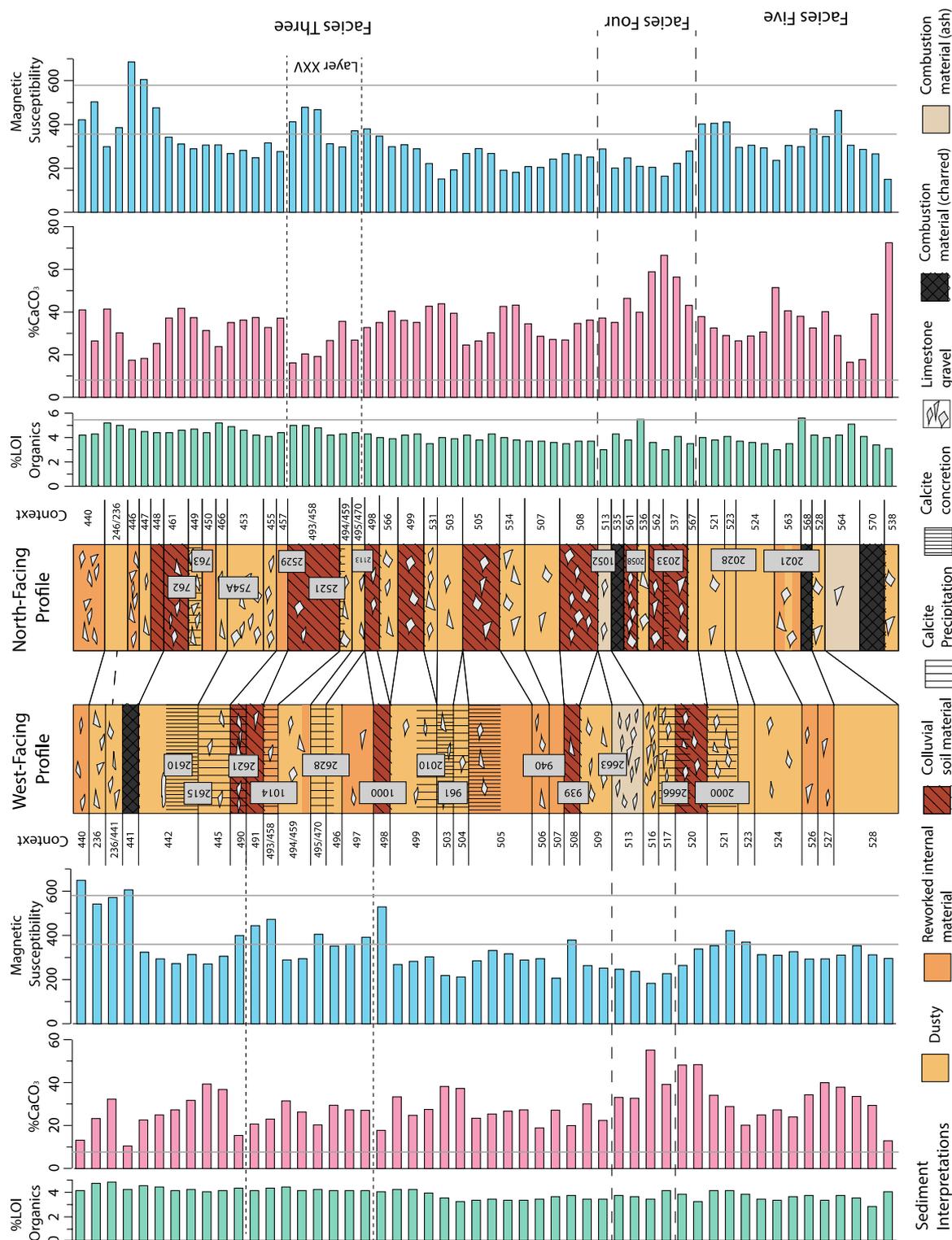


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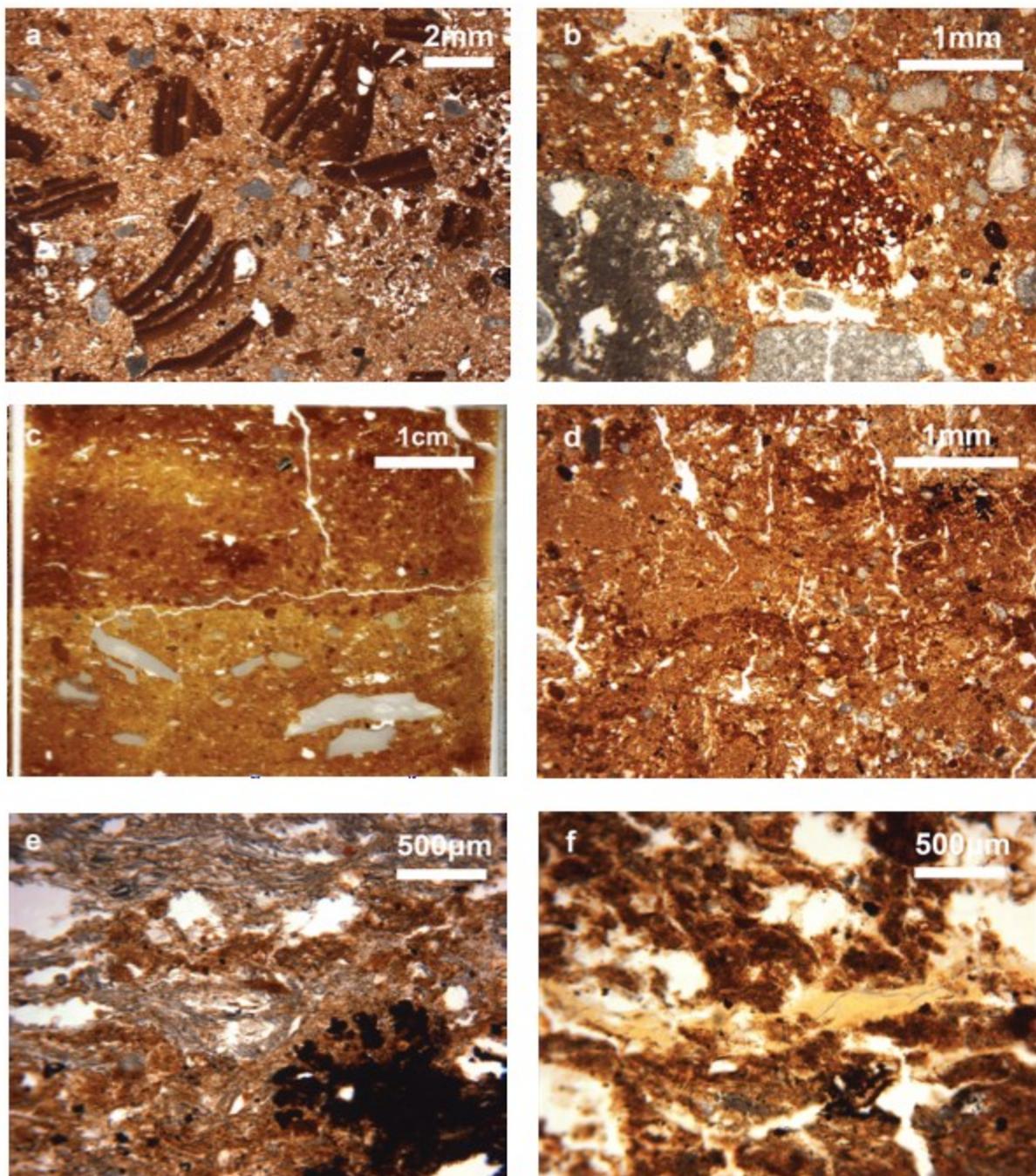


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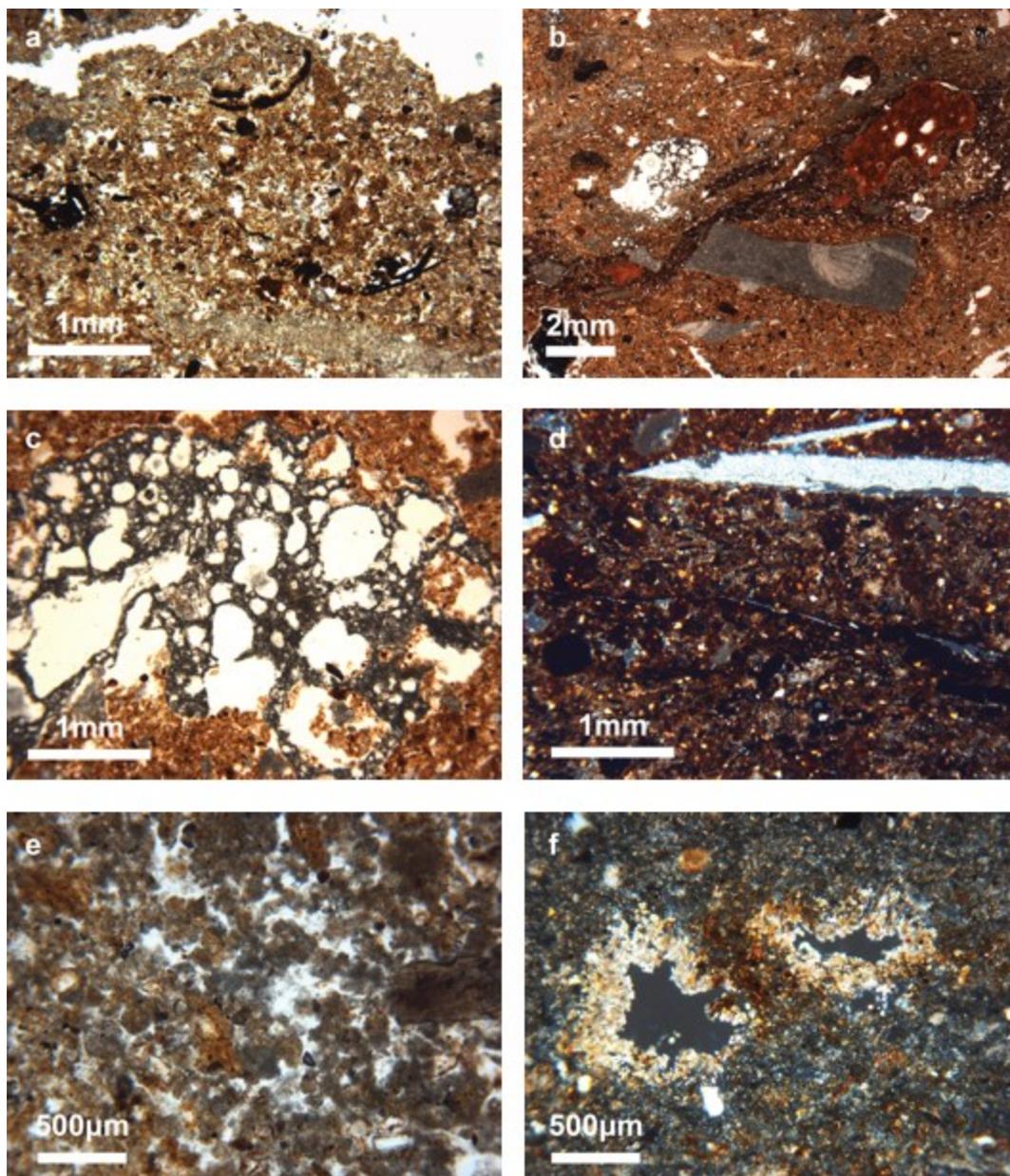


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